

A Procedure to Calibrate Multiparameter Weather Radar Using Properties of the Rain Medium

Eugenio Gorgucci, Gianfranco Scarchilli, and V. Chandrasekar, *Member, IEEE*

Abstract— The joint distribution characteristics of size and shape of raindrops directly translate into features of polarization diversity measurements in rainfall. Theoretical calculations as well as radar observations indicate that the three polarization diversity measurements, namely, reflectivity, differential reflectivity, and specific differential propagation phase, lie in a constrained space that can be approximated by a three-dimensional (3-D) surface. This feature as well as the vertical-looking observation of raindrops are used to determine biases in calibration of the radar system. A simple procedure is developed to obtain the bias in the absolute calibration from polarization diversity observation in rainfall. Simulation study as well as data analysis indicate that calibration errors can be estimated to an accuracy of 1 dB.

Index Terms— Calibration, meteorological radar, polarimetric radar.

I. INTRODUCTION

THE characteristics of polarization diversity measurements in rain are determined by the shape and size distribution of raindrops. The shape of a raindrop is determined by the forces due to the surface tension and hydrostatic and aerodynamic pressures due to airflow around the raindrop. Medium and large raindrops (>2 mm in diameter) are nonspherical. The shape of raindrops changes with the size and can be approximately described by an oblate spheroid with the symmetry axis in the vertical direction [1], [2], [6]. The properties of the size–shape relationship of raindrops translate directly into features of polarization diversity measurements in rainfall. Theoretical calculation and radar observations suggest that polarization diversity measurements of rain, namely, 1) radar reflectivity factor (Z_H), 2) differential reflectivity (Z_{DR}), and 3) specific differential phase (K_{DP}) lie in a limited three-dimensional (3-D) space [3]. This constraint can be utilized to study the calibration errors in a multiparameter weather radar [4].

In principle, it can be stated that the calibration errors in a multiparameter radar can be corrected by enforcing the constraints of observations in rain. However, in practice, the technique to calibrate the multiparameter radar using this self-

consistency procedure is not straightforward and numerous details are important. This paper presents the practical procedure to calibrate multiparameter weather radars. The paper is organized as follows. Section II describes the multiparameter radar measurements in rainfall and their self-consistency. In Section III, the issues involved in calibrating differential reflectivity are described. Section IV presents the procedure to correct errors in absolute calibration. Section V summarizes the results of this paper.

II. SELF-CONSISTENCY OF POLARIZATION DIVERSITY OBSERVATION IN RAIN

Cloud models and measurements of raindrop size distribution (RSD) at the surface and aloft show that a gamma distribution model adequately describes many of natural variations in the RSD [5]

$$N(D) = n_C f_D(D) \quad (\text{m}^{-3} \text{mm}^{-1}) \quad (1)$$

where $N(D)$ is the number of raindrops per unit volume per unit size interval (D to $D + \Delta D$), n_C is the concentration, and $f_D(D)$ is the gamma probability density function (pdf). $f_D(D)$ can be written as

$$f_D(D) = \frac{\Lambda^{\mu+1}}{\Gamma(\mu+1)} e^{-\Lambda D} D^\mu \quad (2)$$

where μ and Λ are parameters of the distribution. The median drop diameter D_0 can be obtained in terms of the parameters of the distribution as

$$D_0 = \frac{3.67 + \mu}{\Lambda} \quad (3)$$

The shape of a raindrop can be approximated by an oblate spheroid with the axis ratio (b/a) given by the empirical relation [6]

$$\frac{b}{a} = 1.03 - 0.062D \quad (4)$$

where D is the equivolumetric spherical diameter of a raindrop in millimeters and a and b are the semimajor and semiminor axis of the raindrop. The radar reflectivity factor of the rain medium can be expressed in terms of the RSD as

$$Z_{H,V} = \frac{\lambda^4}{\pi^5 |K|^2} \int \sigma_{H,V}(D) N(D) dD \quad (\text{mm}^6 \text{m}^{-3}) \quad (5)$$

where $Z_{H,V}$ and $\sigma_{H,V}$ represent the reflectivity factors and radar cross sections at horizontal (H) and vertical (V) polarizations, respectively, λ represents the wavelength, and

Manuscript received December 9, 1996; revised December 16, 1997. This work was supported in part by the National Group for Defence from Hydrogeological Hazards (CNR, Italy), by Mesoscale Alpine Project (MAP-CNR), and by NSF (ATM-9413453).

E. Gorgucci and G. Scarchilli are with the Istituto di Fisica dell'Atmosfera (CNR), Area di Ricerca Roma-Tor Vergata, 00133 Rome, Italy (e-mail: gorgucci@radar.ifa.rm.cnr.it).

V. Chandrasekar is with Colorado State University, Fort Collins, CO 80523 USA.

Publisher Item Identifier S 0196-2892(99)00041-8.

$K = (\epsilon_r - 1)/(\epsilon_r + 2)$, where ϵ_r is the dielectric constant of water. Similarly differential reflectivity (Z_{DR}) and specific differential phase (K_{DP}) can be obtained as

$$Z_{DR} = \frac{\int \sigma_H(D)N(D) dD}{\int \sigma_V(D)N(D) dD} \quad (6)$$

$$K_{DP} = \frac{180\lambda}{\pi} \text{Re} \int [f_H(D) - f_V(D)]N(D) \cdot dD \quad (\text{deg km}^{-1}) \quad (7)$$

where f_H and f_V are the forward scatter amplitudes at H- and V-polarization, respectively. Sarchilli *et al.* [3] have shown that multiparameter radar observations in rain, namely, Z_H , Z_{DR} and K_{DP} , vary over a constrained 3-D space. This constraint enables parameterization of K_{DP} in terms of Z_H and Z_{DR} as

$$K_{DP}^* = CZ_H^\alpha 10^{-\beta Z_{DR}} \quad (8)$$

where K_{DP}^* is the parameterized estimate of K_{DP} , Z_H is in units $\text{mm}^6 \text{m}^{-3}$, and Z_{DR} is in decibels. It should be noted here that K_{DP}^* indicates an estimate based on Z_H and Z_{DR} and not from a profile of differential phase measurements. The coefficients C , α , and β change as a function of the frequency. The values of C , α , and β at S- and C-bands are given as follows.

S-band (10 cm)

$$C = 1.05 \times 10^{-4} \quad \alpha = 0.96 \quad \beta = 0.26. \quad (9a)$$

C-band (5.5 cm)

$$C = 1.46 \times 10^{-4} \quad \alpha = 0.98 \quad \beta = 0.20. \quad (9b)$$

Sarchilli *et al.* [3] have shown that (8) can approximate K_{DP} fairly well in the absence of measurement errors. Thus, we can see that Z_H , Z_{DR} , and K_{DP} exist in a constrained 3-D space and (8) represents a 3-D surface approximation of the variability of Z_H , Z_{DR} , and K_{DP} . In the following section, this self-consistency feature is utilized to study biases in calibration of multiparameter radars.

III. VERTICAL-LOOKING OBSERVATION OF Z_{DR} AND ITS CALIBRATION

The shape of raindrops can be approximated by oblate spheroids. Radar and *in situ* aircraft-based observations show that on average the raindrops are oriented with the symmetry axis in the vertical direction. This implies that the shape of raindrops seen at an elevation angle of 90° is nearly circular. Therefore, Z_{DR} measurements performed with the antenna pointing at an elevation of 90° should be 0 dB. There can be deviations from the 0-dB value of Z_{DR} due to the following reasons:

- 1) mean canting angle of raindrops different from zero;
- 2) ground clutter contribution;
- 3) system bias given by unequal paths or gains in the H- and V-polarization channels.

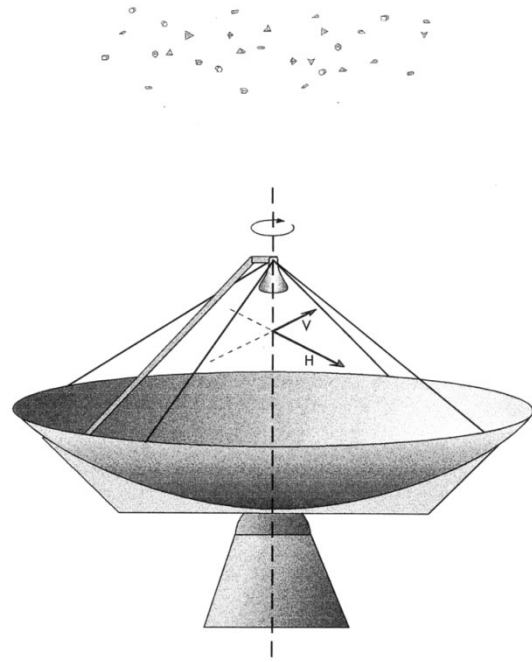


Fig. 1. Variation of H- and V-polarization directions due to the rotation of the radar antenna about the vertical.

Items 1) and 2) depend on H- and V-polarization directions (see Fig. 1). The average Z_{DR} taken over all possible orientations of H- and V-polarizations, with the antenna looking vertical for items 1) and 2), go to zero. However if there is nonzero Z_{DR} due to the system bias, it does not change with the different H and V orientations looking vertical. In many radars, the different H and V orientations can be achieved by changing the azimuth positioning over zero to 360° , keeping the elevation angle at 90° . However, the exact procedure depends on the set up of the scanning servo system of the radar. In summary, the average Z_{DR} computed with all possible orientations of the polarization states of the radar antenna pointing in the vertical direction should be zero. Any nonzero observation can be directly attributed to the radar system bias between the two polarization channels. In the following, we show the estimation of the system Z_{DR} bias using vertical-looking data from the NCAR, CP-2 radar, taken over the Kennedy Space Center (KSC), FL, during the Convection and Precipitation/Electrification Experiment (CaPE) [7]. CaPE was conducted in the Central Florida region during the summer of 1991.

Fig. 2 shows the azimuthal profile of Z_{DR} obtained during the rotation of radar antenna about the vertical; each value of Z_{DR} represents an average of 2 km in altitude and 5° in azimuth. When the azimuth angle in horizontal axis exceeds 360° , the exact position of the antenna can be obtained subtracting multiples of 360° from the indicated angle in Fig. 2. We can see from Fig. 2 that the measured Z_{DR} has a periodic structure in the azimuth, with a periodicity of 180° . The mean level of the Z_{DR} value is about 0.15 dB. The full excursion of the Z_{DR} profile as a function of the azimuth is 0.3 dB. This excursion is due to the variability in the ground clutter response with the rotation of the antenna. Vertical-looking

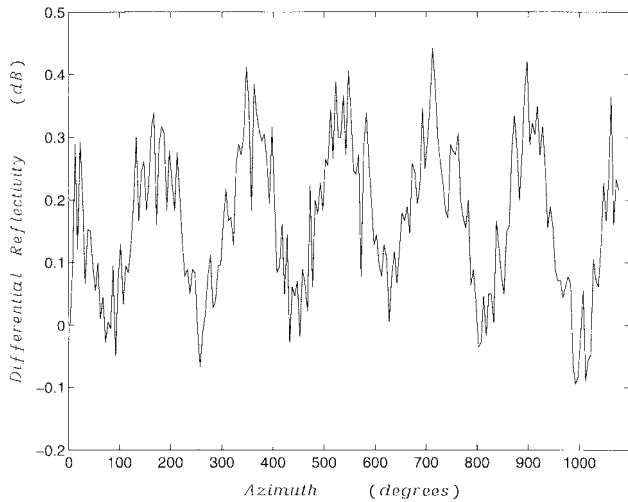


Fig. 2. Profile of vertical-looking Z_{DR} measurements, collected in rain medium below 3.5 km in altitude, as a function of the azimuth for the radar antenna in continuous rotation about the vertical.

radar observations were made for several minutes, rotating continuously about the vertical. Based on the observations from Fig. 2, we can see that it is important to average Z_{DR} over full cycles of 360° in azimuth to avoid bias due to the direction-dependent response. The Z_{DR} profile was obtained from data below 3.5 km in altitude to stay in rain medium. The same profile was observed at various altitudes, and the estimated mean does not change as a function of height, again confirming that the mean Z_{DR} is the system bias. The standard deviation of Z_{DR} evaluated from the data was 0.3 dB. The total number of points used in the mean estimates was about 10000. Therefore, the standard deviation in the estimate of the Z_{DR} bias is 0.003 dB.

IV. ABSOLUTE CALIBRATION

The absolute calibration of the radar is important for quantitative application of radar measurements, such as rainfall rate estimation. The absolute calibration of the radar depends on several factors, such as the antenna gain, gain of amplifiers, loss in waveguides, etc. In this section, we describe a procedure for using the properties of the rain medium to obtain absolute calibration of the radar. The backscattered mean power received by the radar (P_r) from a resolution volume can be written as

$$P_r = C \frac{Z}{R^2} \quad (10)$$

where Z is the reflectivity factor, R is the range to the volume, and C is the radar constant. Accurate knowledge of C is very important for estimating Z from the received power. Typically the radar constant is known within couple of decibels from past calibration. This calibration may drift by a few decibels due to changes in the properties of the various amplifiers and mixers along the line. Invoking the principle of self-consistency, we can estimate the drift in the calibration of the radar constant. In the absence of measurement errors, a scatter diagram of measured K_{DP} and K_{DP}^* obtained from (8) should lie about a 1:1 line. However, in the presence of a bias in the absolute

gain of the radar system, a scattergram of K_{DP} versus K_{DP}^* will deviate from a 1:1 line. The slope s of this scattergram depends on the amount of bias in the absolute calibration, and it can be obtained as a simple expression based on (8) as

$$\text{Bias} = 10.4 \times \log(s) \quad (\text{dB}). \quad (11)$$

The standard deviation of the bias can be approximated by

$$\sigma(\text{Bias}) = 4.52 \frac{\sigma(s)}{E(s)} \quad (\text{dB}) \quad (12)$$

where $\sigma(s)$ and $E(s)$ are the standard deviation and the mean value of the slope, respectively.

In the following sections, we present theoretical studies as well as simulations that evaluate the accuracy of the calibration procedure.

A. Estimation of the Bias in Calibration

The multiparameter measurements in rainfall, namely, Z_H , Z_{DR} , and K_{DP} , lie in a constrained 3-D space that can be approximated by a surface given by (8). K_{DP} is the slope of the differential phase profile with range and is independent of the calibration of the radar. Z_{DR} is a relative power measurement and can be calibrated fairly accurately, as shown in the previous section. Any bias in the absolute calibration of the radar (error in the assumed radar constant) will result in a biased value of reflectivity. This bias will also produce a discrepancy in the self-consistency of the triplet of multiparameter radar measurements Z_H , Z_{DR} , and K_{DP} .

The principle for obtaining the bias in calibration is fairly straightforward. However, numerous details are important to obtain the bias. Firstly, the measurement error and nature of Z_H , Z_{DR} , and K_{DP} are very different. Z_H and Z_{DR} are power measurements, whereas K_{DP} is derived from Φ_{DP} , which is the differential phase measurement. Z_H and Z_{DR} are point measurements, whereas K_{DP} is evaluated over a propagation path. A mean K_{DP} value is evaluated typically over a path of few kilometers in length. The measurement error in K_{DP} depends on the number of the range gates along the path, the range resolution, and the profile of Φ_{DP} . In this section, we analyze various profiles of Z_H , Z_{DR} , and Φ_{DP} to study the implication on the estimation of bias in calibration. Sarchilli *et al.* [3] have shown that parameterized Φ_{DP}^* profile can be constructed from measurement of Z_H and Z_{DR} using (8). Any error in absolute calibration will result in a deviation from the measured profile of Φ_{DP} , where the difference between Φ_{DP} and Φ_{DP}^* will progressively increase with range. The actual Φ_{DP} profile depends on the Φ_{DP} value at the starting range bin (Φ_{DP}^o), which is the differential phase due to radar system. Therefore, if Φ_{DP}^o is not known accurately, there will be a constant shift between Φ_{DP} and Φ_{DP}^* profiles, even in the absence of any calibration bias. Φ_{DP}^o can vary over several degrees within a day, as shown in the histogram of Fig. 3, where it can be seen that Φ_{DP}^o varies by 10° . We note here that this spread of 10° includes spread due to measurement errors [8]. Any technique that matches profiles of Φ_{DP} and Φ_{DP}^* will be affected by uncertainty in the knowledge of initial Φ_{DP}^o . Therefore, to overcome this

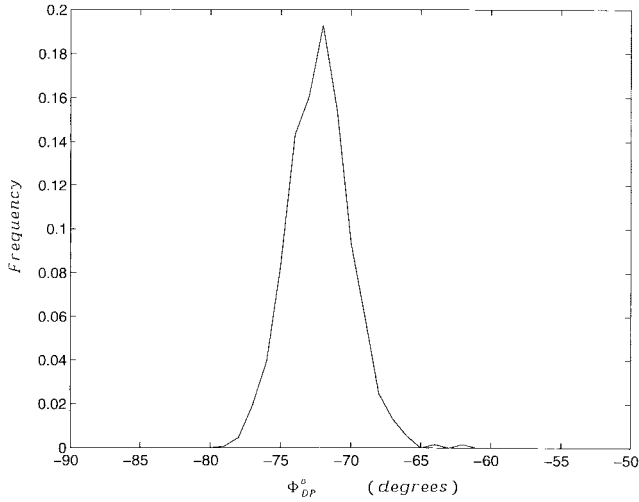


Fig. 3. Histogram of the differential profile phase shift Φ_{DP}^o at the starting range bin.

problem, the range variation of Φ_{DP} and Φ_{DP}^* profiles is utilized in our analysis.

Finally, we want to point out that important physical factors, such as attenuation and hail contamination, can affect the accuracy of the method. For the calibration purpose, we should avoid using hail-contaminated data because the procedure is strictly for rain medium. As for the issue of attenuation, we can note that, according to (8), the bias B_H in Z_H is equivalent to the bias B_{DR} in Z_{DR} as

$$B_{DR} = 0.1 \frac{\alpha}{\beta} B_H \quad (\text{dB}). \quad (13)$$

Thus, to ensure the accuracy of Z_H of 1 dB, we need to have B_{DR} below 0.4 dB at S -band and 0.5 dB at C -band. The bias B_{DR} due to differential attenuation is proportional to the total differential phase and can be approximated as

$$B_{DR} = b\Phi_{DP} \quad (\text{dB}) \quad (14)$$

where b is 0.0033 and 0.014 dB/km at S - and C -band, respectively [9]–[10]. Combining (13) and (14), we can obtain the cumulative differential phase corresponding to the Z_{DR} bias due to differential attenuation. Using this procedure, we can obtain that when cumulative differential phase at S -band reaches 112° , then the bias in Z_{DR} due to differential attenuation appears equivalent to 1-dB error in Z_H . Similar Φ_{DP} value at C -band is 35° . Thus, we see that Z_{DR} values have to be corrected for differential attenuation at C -band, whereas at S -band, it is needed when we use data over a long path, resulting in cumulative Φ_{DP} of over 112° .

B. Error Structure of K_{DP} and Calibration Bias Estimates

The knowledge of the error structure of K_{DP} , K_{DP}^* and bias is critical to the understanding of accuracy of the calibration procedure. Theoretical estimates of the standard errors in K_{DP} , K_{DP}^* and bias in calibration are obtained in this section for a uniform precipitation path.

When the reflectivity factor is uniform along the rainfall path, we can estimate the mean value of the specific differential

phase shift K_{DP} through the least-squares fit procedure as

$$\hat{K}_{DP} = \frac{\sum_{i=1}^N [(\Phi_{DP})_i - \bar{\Phi}_{DP}](r_i - \bar{r})}{2 \sum_{i=1}^N (r_i - \bar{r})^2} \quad (15)$$

where \hat{K}_{DP} is the estimate of K_{DP} , $(\Phi_{DP})_i$ is the two-way cumulative differential phase shift at the range bin i , $\bar{\Phi}_{DP}$ is the mean value of Φ_{DP} along the path, r_i is the distance of the range bin from the radar, \bar{r} is the average value of the path length, and N is the total number of range bins. Assuming $r_i = i\Delta r$, where Δr is the range resolution it can be easily shown that

$$\bar{r} = \left(\frac{N+1}{2} \right) \Delta r \quad (16a)$$

$$\sum_{i=1}^N (r_i - \bar{r})^2 = \left[\frac{N(N-1)(N+1)}{12} \right] \Delta r^2. \quad (16b)$$

Substituting (16) in (15), we get

$$\hat{K}_{DP} = \frac{\sum_{i=1}^N [(\Phi_{DP})_i - \bar{\Phi}_{DP}] \left(i - \frac{N+1}{2} \right) \Delta r}{2 \frac{N(N-1)(N+1)}{12} \Delta r}. \quad (17)$$

Subsequently, the variance of \hat{K}_{DP} can be obtained as

$$\begin{aligned} \text{var}(\hat{K}_{DP}) &= \text{var}(\Phi_{DP}) \frac{\sum_{i=1}^N \left[\left(i - \frac{N+1}{2} \right) \Delta r \right]^2}{4 \left[\frac{N(N-1)(N+1)}{12} \Delta r^2 \right]^2} \\ &= \frac{\text{var}(\Phi_{DP})}{4 \frac{N(N-1)(N+1)}{12} \Delta r^2}. \end{aligned} \quad (18)$$

We can note here that (18) converges to (5), given by Ryzhkov and Zrnice [11], for large N . It can be observed from (18) that the variance of \hat{K}_{DP} decreases by increasing the number of range bins, utilized in the estimate of K_{DP} , and the radial resolution Δr .

Following a similar procedure as above, we can estimate the mean value of K_{DP}^* through the least-squares fit as

$$\hat{K}_{DP}^* = \frac{\sum_{i=1}^N [(\Phi_{DP}^*)_i - \bar{\Phi}_{DP}^*] \left(i - \frac{N+1}{2} \right) \Delta r}{2 \sum_{i=1}^N \left[\left(i - \frac{N+1}{2} \right) \Delta r \right]^2} \quad (19)$$

where the $(\Phi_{DP}^*)_i$ can be written

$$(\Phi_{DP}^*)_i = 2(K_{DP}^*)_i \Delta r. \quad (20)$$

With simple algebraic manipulations, (19) can be reduced to

$$\hat{K}_{DP}^* = \frac{\sum_{i=1}^N (K_{DP}^*)_i \left[\left(i - \frac{N+1}{2} \right) \Delta r \right]^2}{\sum_{i=1}^N \left[\left(i - \frac{N+1}{2} \right) \Delta r \right]^2}. \quad (21)$$

The variance of the estimate (6) becomes [12]

$$\text{var}(\hat{K}_{DP}^*) = \frac{\sum_{i=1}^N \text{var}(K_{DP}^*)_i \left[\left(i - \frac{N+1}{2} \right) \Delta r \right]^4}{\left[\sum_{i=1}^N \left(i - \frac{N+1}{2} \right)^2 \Delta r^2 \right]^2}. \quad (22)$$

In each range bin, $\text{var}(K_{DP}^*)$ is the same and can be written [3]

$$\text{var}(K_{DP}^*) = (K_{DP}^*)^2 \nu \quad (23)$$

where

$$\nu = [0.053\alpha^2 \text{var}(10\log Z_H) + 5.3\beta^2 \text{var}(Z_{DR})]. \quad (24)$$

Then (22) can be expressed as

$$\text{var}(\hat{K}_{DP}^*) = (K_{DP}^*)^2 \nu \frac{\sum_{i=1}^N \left[\left(i - \frac{N+1}{2} \right) \Delta r \right]^4}{\left[\sum_{i=1}^N \left(i - \frac{N+1}{2} \right)^2 \Delta r^2 \right]^2}. \quad (25)$$

We want to point out that, unlike the variance of \hat{K}_{DP} , the variance of \hat{K}_{DP}^* is independent of the range resolution Δr ; moreover, it decreases slowly by increasing the number of the averaged range bins.

The parameters assumed for the computation of the radar measurement error are as follows:

- 1) radar frequency 3 GHz;
- 2) pulse repetition time 1 ms;
- 3) number of sample pairs 64;
- 4) Doppler spectrum width 2 m/s;
- 5) zero lag cross-correlation coefficient 0.99;
- 6) range resolution 0.3 km.

With those assumptions, we obtain the standard deviations in the estimates of Z_H , Z_{DR} , and Φ_{DP} to be 0.7 dB, 0.15 dB, and 1° , respectively.

Fig. 4 shows the standard deviation of the estimates \hat{K}_{DP}^* and \hat{K}_{DP} for different values of Z_H . The Z_{DR} value that corresponds to each Z_H shown in Fig. 4 was chosen to be the average value of a Z_H, Z_{DR} scatterplot in rainfall [13]. The Z_{DR} values corresponding to the reflectivity values 30, 35, 40, 45, 50, and 53 dBZ shown in Fig. 4 are 0.9, 0.95, 1.3, 1.6, 1.9, and 2.0 dB, respectively. We need to note here that the standard deviation of \hat{K}_{DP} does not change with reflectivity. Figs. 5 and 6 show the fractional standard error (FSE) in the estimates \hat{K}_{DP} and \hat{K}_{DP}^* as a function of number of range bins over which the least-squares estimates are computed. It should be noted here that the various curves in Fig. 5 correspond to the same Z_H, Z_{DR} pair shown in Fig. 4.

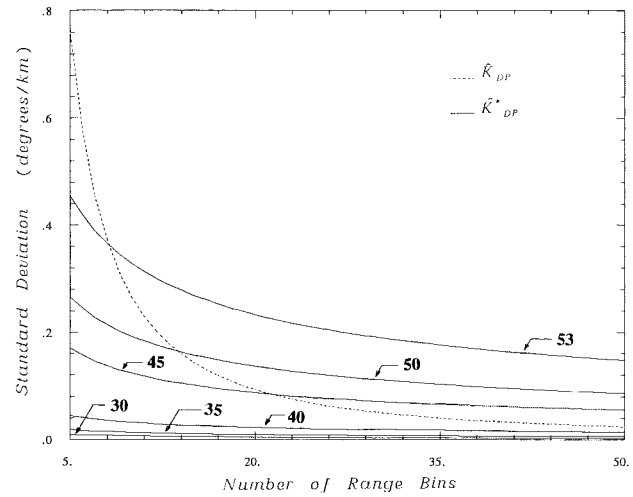


Fig. 4. Standard deviation of the estimate \hat{K}_{DP}^* for different values of Z_H (continuous lines), and of the estimate \hat{K}_{DP} (dashed line) as a function of number of range bins.

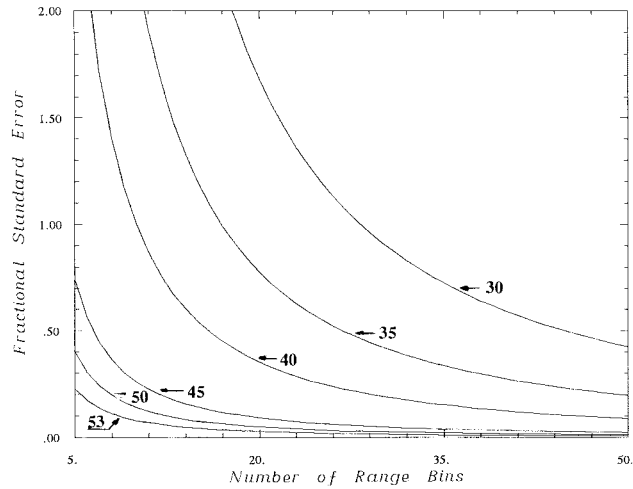


Fig. 5. Fractional standard error of the estimate \hat{K}_{DP} for different values of Z_H as a function of number of range bins.

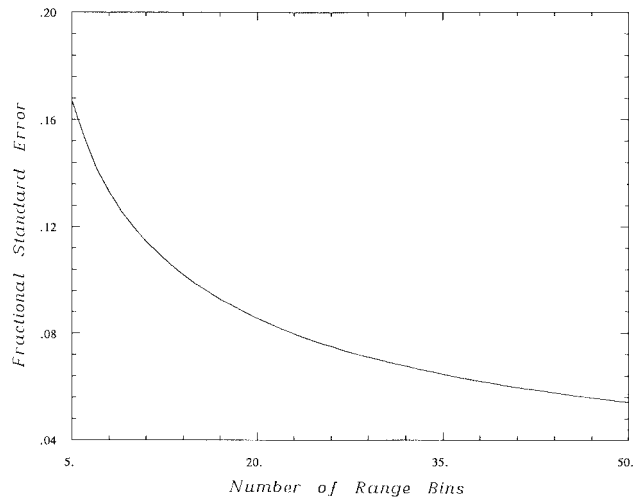


Fig. 6. Fractional standard error of the estimate \hat{K}_{DP}^* as a function of number of range bins.

The slope s of the scatter diagram between \hat{K}_{DP}^* and \hat{K}_{DP} is utilized to estimate the bias in calibration. The least-squares estimate of the slope of a scattergram between the estimates \hat{K}_{DP}^* and \hat{K}_{DP} is given by

$$\hat{s} = \frac{\sum_{i=1}^M (\hat{K}_{DP}^*)_i (\hat{K}_{DP})_i}{\sum_{i=1}^M [(\hat{K}_{DP}^*)_i]^2} \quad (26)$$

where M is the number of the estimates. The variance in the estimate of the slope can be written as

$$\text{var}(\hat{s}) = \frac{1}{M} \frac{\text{var}(\hat{K}_{DP}^*)}{\text{var}(\hat{K}_{DP})} \quad (27)$$

where ρ is the correlation between the estimates of K_{DP}^* and K_{DP} [14]. The value of ρ is very close to unity, typically of the order of 0.99, as shown through simulation in the next section. It can be seen from Fig. 4 that the ratio of the standard deviation of the two K_{DP} estimates commonly takes values between 0.1 and two. Therefore, the standard deviation of \hat{s} can be approximated to be $0.3/\sqrt{M}$. Thus, if we take about 100 estimates of K_{DP}^* and K_{DP} in moderate to heavy rain, the bias can be estimated to an accuracy of about 0.1 dB. We need to note here that this estimate of accuracy in bias will vary depending on the accuracy of the two K_{DP} estimates, which in turn depend on the path length over which K_{DP}^* and K_{DP} are estimated. Nevertheless, this first approximation provides an idea about the accuracy of the slope estimate.

C. Simulation Study of Rain Profiles

The various range profiles of Φ_{DP} , Z_H , and Z_{DR} are constructed by simulation with the following characteristics:

- 1) total path length of 90 km;
- 2) range resolution 300 m;
- 3) gradients of reflectivity in range vary between 0.5 and 50 dB/km;
- 4) maximum reflectivity along the path is less than 55 dBZ.

The reflectivity profile of the precipitation path is constructed such that, when the reflectivity is equal to 55 dBZ or drops below 0 dBZ, the direction of the gradient is changed. Once the reflectivity is fixed at a range resolution volume, the parameters of the RSD, namely, N_0 , D_0 , and μ , are chosen randomly under the constraint that the RSD yields the current reflectivity value in the range resolution cell. Subsequently, the values of Z_{DR} and K_{DP} are computed for the range location. This procedure yields a wide variety of range profiles of Z_H , Z_{DR} , and K_{DP} so that the technique to evaluate calibration bias can be tested under all conditions. The value of $\Phi_{DP}(r)$ over range is computed from K_{DP} as $2\sigma K_{DP}(r)\Delta r$, where Δr is the range resolution. Once the mean values of Z_H , Z_{DR} , and Φ_{DP} are fixed at each range resolution cell, the radar estimates with measurement error are simulated using the procedure given by Chandrasekar *et al.* [15].

Fig. 7 shows sample profiles of Z_H , Φ_{DP} , and Φ_{DP}^* that are used in the simulation. There are several ways to compare the true and parameterized values of K_{DP} . However, practical constraints, such as error structure and range cumulative nature

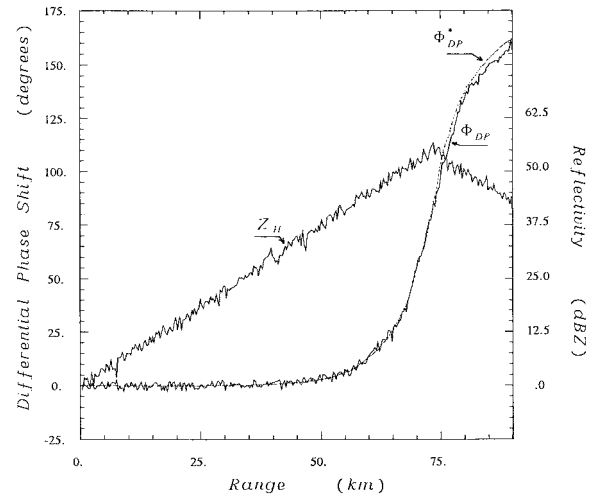


Fig. 7. Sample profiles of Z_H , Φ_{DP} , and Φ_{DP}^* obtained in simulation as a function of range.

of differential propagation phase measurements, restrict the possibilities. In our analysis, K_{DP} is estimated as the slope of the range profile of Φ_{DP} to minimize its standard error. This procedure has advantages and disadvantages. Using a least-squares fit to the Φ_{DP} profile, we can estimate K_{DP} accurately in the case of a fairly long path (at least 15–20 range gates). However, when the path over which the least-squares fit is done becomes long, K_{DP} cannot be uniform anymore within the path and the estimates are affected by gradients. When there are gradients in the path over which K_{DP} is estimated, a bias is introduced in the K_{DP} estimate [16]. In the context of estimating bias in the calibration of the radar system, even small errors in the estimate of K_{DP} can be significant. To circumvent this problem, we apply an identical least-squares estimate to the parameterized Φ_{DP}^* profile. Therefore, an average K_{DP}^* estimate over the path is obtained performing a least-squares fit to the Φ_{DP}^* profile. This procedure is adopted so that if there is a bias in the estimate of K_{DP} due to gradients in the path, the same bias will be there in the estimate of K_{DP}^* . Thus, any difference between K_{DP} and K_{DP}^* will be due only to the bias in absolute calibration of the radar. Fig. 8 shows the scatterplot between the estimates of K_{DP} and K_{DP}^* obtained as slopes of Φ_{DP} and Φ_{DP}^* profiles, respectively. The various points correspond to the different paths considered in our simulation study. We can see that the scatter between the two estimates lies about at 45° . The slope of the scatter is 0.996, and the standard deviation in the estimate of the slope is 0.007. This slope corresponds to a bias of 0.02 dB. We can see that the accuracy of slope and bias estimates obtained from simulation are in the approximate range given theoretically from (27). Fig. 9 shows a similar scatterplot, except that the simulation includes a calibration bias of 1 dB. The estimated slope of the scatter in Fig. 9 is 0.7926, corresponding to a bias of -1.05 dB. The accuracy of the bias estimate depends on several factors, such as accuracy in measurements of Φ_{DP} , Z_H , and Z_{DR} , path length in estimation of K_{DP} , and the number of data points in the scattergram of \hat{K}_{DP}^* and \hat{K}_{DP} estimates. Based on our simulations, it appears that, if we have sufficient data points

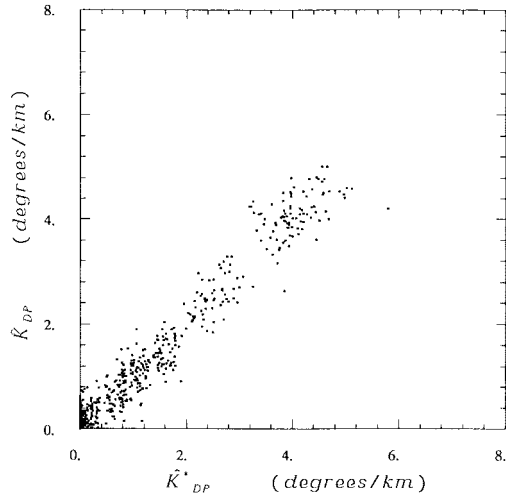


Fig. 8. Scatterplot between the estimates \hat{K}_{DP} and \hat{K}_{DP}^* obtained as slopes of Φ_{DP} and Φ_{DP}^* , respectively.

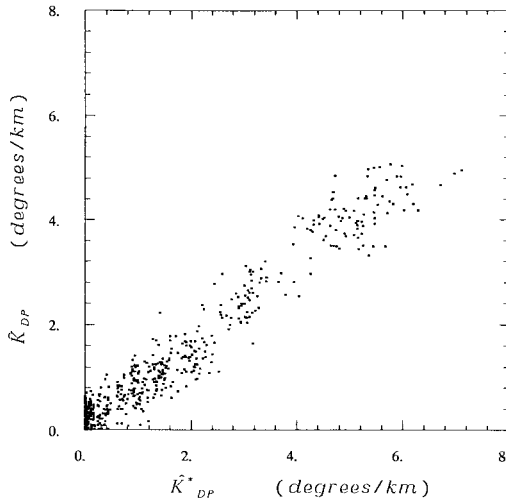


Fig. 9. Same as Fig. 9, introducing a bias of 1 dB in simulated Z_H measurements.

for at least 100 estimates of K_{DP} , we can estimate the bias in the absolute calibration to an accuracy of a few tenth of 1 dB. One important point to be noted here is the success of this procedure depends on obtaining many accurate estimates of K_{DP} over various paths. Therefore, data from light rain (where there is not significant K_{DP}) should not be used in the calibration data set. A large amount of data from moderate to heavy rain ($>50 \text{ mm h}^{-1}$) will be very useful.

D. Data Analysis

Data collected during the CaPE experiment is used to validate the calibration procedure developed in this paper. During the later part of the CaPE experiment, the NCAR CP-2 radar was equipped with an auxiliary signal processor developed at Colorado State University to obtain differential phase measurements in addition to Z_H and Z_{DR} . Table I shows the main features of the CP-2 radar that are relevant to this paper. Fig. 10 shows a typical range profile of radar data used in the analysis. Fig. 11 shows the scatterplot between \hat{K}_{DP} and \hat{K}_{DP}^* estimates for data collected on September 29, 1991,

TABLE I
CHARACTERISTICS OF THE CP-2 RADAR

Characteristic	CP-2 (S-Band)
Polarization Type	linear
Wavelength (cm)	10.7
Peak Power (kW)	1200
Pulse Length (μs)	1.0
PRF (s^{-1})	960
Antenna Type	Center fed paraboloid
Antenna Size (m)	8.5
Beamwidth (degrees)	0.93
Polarizations transmitted	Linear V or H
Polarizations received	Copolar to transmit
Maximum Sidelobe level (dB)	-21
Polarization Control Method	Ferrite switch
Polarization Control Period	Pulse by pulse

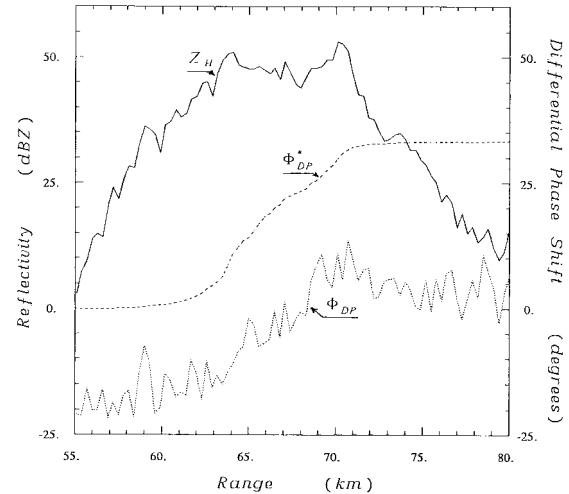


Fig. 10. Typical range profiles of Z_H , Φ_{DP} , and Φ_{DP}^* obtained from radar data collected on September 29, 1991, by CP-2 during the CaPE experiment.

during the CaPE experiment. The measurements were obtained integrating 64 sample pairs, with a typical scan rate of 8 deg s^{-1} . The \hat{K}_{DP} estimates were obtained over 30 range gates. The scatterplot between \hat{K}_{DP} and \hat{K}_{DP}^* estimates has a deviation away from 45° line that corresponds to a bias of -1.1 dB . Data collected over a long time during the day have shown the stability of the results. This bias determined from data was eventually traced to changes in system component characteristics in the radar.

V. SUMMARY AND CONCLUSION

The properties of shape and size distribution of raindrops translate into features of polarization diversity measurement. Polarization diversity observations in rain, namely, reflectivity, differential reflectivity, and specific differential propagation phase, lie in a constrained 3-D space. This self-consistency feature has been used to evaluate biases in the absolute calibration of the radar. In addition, vertical-looking observations in rainfall with antenna rotating about the vertical axis have been used to calibrate Z_{DR} measurements. The self-consistency feature of polarization diversity measurements are

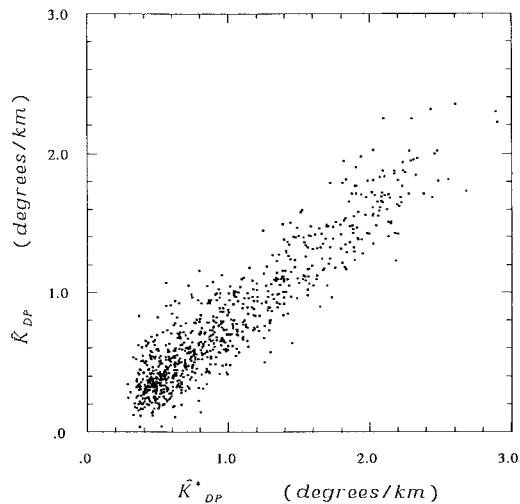


Fig. 11. Scatterplot between the estimates \hat{K}_{DP} and \hat{K}_{DP}^* obtained as slopes of Φ_{DP} and Φ_{DP}^* , respectively. Radar data were collected on September 29 1991, by CP-2 during the CaPE experiment.

used to parameterize K_{DP} as a function of reflectivity and differential reflectivity. Subsequently, this parameterized K_{DP}^* is compared against directly measured K_{DP} to evaluate the bias in the absolute calibration of the radar. Multiparameter radar measurements over a large number of precipitation paths were simulated to evaluate the feasibility of the procedure developed in this paper to estimate the bias in absolute calibration. Theoretical analysis as well as simulation study were used to demonstrate the sensitivity of the slope of scattergram between the estimates of K_{DP}^* (involving self-consistency principle) and measured K_{DP} to biases in absolute calibration. A simple expression is provided to estimate the absolute calibration bias from the slope of the scattergram between \hat{K}_{DP}^* and \hat{K}_{DP} estimates.

The calibration procedure developed in this paper was applied to multiparameter radar data from CaPE experiment to evaluate the radar system bias. Data analysis and simulation study presented in this paper show that absolute calibration biases can be estimated to an accuracy less than 1 dB. The actual accuracy depends on several factors, such as accuracies of Z_H , Z_{DR} , and Φ_{DP} measurements, path lengths over which K_{DP} is computed from Φ_{DP} , as well as the number of data points used in the estimates of K_{DP}^* and K_{DP} . One important point to note here is that, to obtain good estimates of bias in absolute calibration, significant Φ_{DP} accumulation over the range profile has to be utilized. Data from light rain will not be sufficient to achieve this goal. Extensive data sets from moderate to heavy rainfall ($>50 \text{ mm h}^{-1}$) are needed to apply this procedure successfully.

ACKNOWLEDGMENT

The authors are grateful to Mr. P. Iacovelli for drafting and typing.

REFERENCES

[1] T. A. Seliga, and V. N. Bringi, "Potential use of the reflectivity at orthogonal polarizations for measuring precipitation," *J. Appl. Meteorol.*, vol. 15, pp. 69–76, 1976.

- [2] V. Chandrasekar, W. A. Cooper, and V. N. Bringi, "Axis ratios and oscillation of raindrops," *J. Atmos. Sci.*, vol. 45, pp. 1325–1333, 1988.
- [3] G. Scarchilli, E. Gorgucci, V. Chandrasekar, and A. Dobaie, "Self-consistency of polarization diversity measurement of rainfall," *IEEE Trans. Geosci. Remote Sensing*, vol. 34, pp. 22–26, Jan. 1996.
- [4] V. Chandrasekar, V. N. Bringi, N. Balakrishnan, and D. S. Zrnica, "Error structure of multiparameter radar and surface measurements of precipitation, Part III: Propagation differential phase shift," *J. Atmos. Ocean. Technol.*, vol. 7, pp. 621–629, 1990.
- [5] C. W. Ulbrich, "Natural variations in the analytical form of raindrop size distributions," *J. Clim. Appl. Meteorol.*, vol. 22, pp. 1764–1775, 1983.
- [6] H. R. Pruppacher and R. L. Pitter, "A semi-empirical determination of the shape of cloud and raindrops," *J. Atmos. Sci.*, vol. 28, pp. 86–94, 1971.
- [7] G. B. Foote, "Scientific overview and operations plan," Nat. Center Atmos. Res., Boulder, CO, 1991, p. 145.
- [8] A. B. Kostinski, "Fluctuations of differential phase and radar measurements of precipitation," *J. Appl. Meteorol.*, vol. 33, pp. 1176–1181, 1994.
- [9] V. N. Bringi, V. Chandrasekar, N. Balakrishnan, and D. S. Zrnica, "An examination of propagation effects in rainfall on radar measurements at microwave frequencies," *J. Atmos. Ocean. Technol.*, vol. 7, pp. 829–840, 1990.
- [10] G. Scarchilli, E. Gorgucci, V. Chandrasekar, and T. A. Seliga, "Rainfall estimation using polarimetric techniques at C-band frequencies," *J. Appl. Meteorol.*, vol. 32, pp. 1150–1160, 1993.
- [11] A. Ryzhkov and D. S. Zrnica, "Assessment of rainfall measurement that uses specific differential phase," *J. Appl. Meteorol.*, vol. 35, pp. 2080–2090, 1996.
- [12] A. Papoulis, *Probability, Random Variables, and Stochastic Processes*. New York: McGraw-Hill, 1965, p. 583.
- [13] V. N. Bringi, V. Chandrasekar, P. Meischner, J. Hubbert, and Y. Golestani, "Polarimetric radar signatures of precipitation at S- and C-bands," *Proc. Inst. Elect. Eng. F*, vol. 138, no. 2, pp. 109–115, 1991.
- [14] M. G. Kendall and A. Stuart, *The Advanced Theory of Statistics*, vol. 2. New York: Griffin, 1973, p. 723.
- [15] V. Chandrasekar, V. N. Bringi, and P. J. Brockwell, "Statistical properties of dual polarized radar signals," preprints *Proc. 23th Conf. Radar Meteorol.* Snowmass, CO: Amer. Meteorol. Soc., 1986, pp. 154–157.
- [16] G. Scarchilli, E. Gorgucci, and V. Chandrasekar, "Influence of reflectivity on Φ_{DP} and K_{DP} estimate," preprints *Proc. 26th Conf. Radar Meteorol.* Norman, OK: Amer. Meteorol. Soc., 1993, pp. 103–105.



Eugenio Gorgucci received the doctoral degree in physics from La Sapienza University, Rome, Italy.

He has been with the Institute of Atmospheric Physics (IFA), National Research Council (CNR), Rome, as a Research Scientist since 1976. His research interests include radar meteorology, signal processing, wave propagation, and scattering.



Gianfranco Scarchilli received the doctoral degree in electronics engineering from La Sapienza University, Rome, Italy.

He is with the Institute of Atmospheric Physics (IFA) of the National Research Council (CNR), Rome, where he is a Research Scientist. His research interests include radar meteorology, signal processing, and wave propagation.

V. Chandrasekar (S'83–M'83), photograph and biography not available at the time of publication.