

Simulation of nitrogen behaviour of soil-plant systems

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4.7 Use of tracers and computer simulation techniques to assess mineralization and immobilization of soil nitrogen

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1. NAME OF MODEL

TRAMIN - Use of tracer and computer simulation techniques to assess mineralization and immobilization of soil nitrogen.

2. SYSTEM MODELED

The model describes the mineralization and immobilization of soil nitrogen during microbial growth and decay under laboratory conditions.

3. OBJECTIVES

The objectives of this investigation were to develop and test concepts of microbial growth and decay in soil and to study the mineralization and immobilization of soil N using tracer and computer simulation techniques.

4. TIME SCALE

The programme was designed to simulate the duration of the laboratory experiments which lasted from 42 to 84 days. The time resolution was fixed at 0.01 day.

5. DIAGRAM

The flow chart for transfer of C and N involved in the mineralization and immobilization processes is presented in Fig. 1. Four versions of the model are available depending on whether tracer ^{15}N and/or ^{14}C were used. In the basic version, C and N behavior in the mineralization-immobilization processes were simulated. In the other versions, the number of N and/or C integrals were doubled when ^{15}N and/or ^{14}C tracers were used.

6. LEVELS

See 7.

7. GOVERNING EQUATIONS

In models that describe soil organic matter turnover, the microbial biomass transforms C substrates in soil (Paul and Van Veen, 1978) and is directly involved in the N mineralization-immobilization processes. Microbial growth is primarily limited by substrate and nutrient availability and by the physical conditions prevailing in soil. Soil microorganisms grow on a variety of C substrates. For simulation purposes, the exogenous sub-

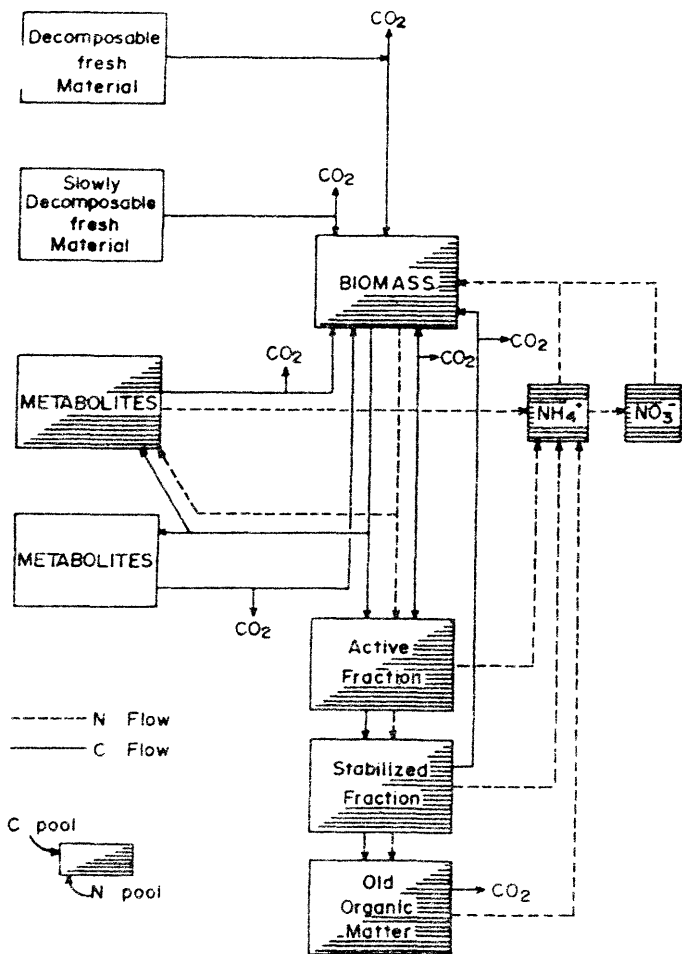


Fig. 1. Flow chart for transfer of C and N involved in the mineralization and immobilization processes.

strates added to soil can be adequately divided into the following categories: (1) decomposable fresh materials consisting of simple carbohydrates; (2) slowly decomposable fresh materials such as cellulose and hemicellulose; (3) materials containing C and N such as proteins, amino sugars, and other nitrogenous compounds; and (4) complex materials such as lignin. In the model, the soil C and N substrates were divided into active-(C+N), stabilized-(C+N) and old-(C+N) (Fig. 1). The techniques and assumptions used to determine the pool sizes are described elsewhere (Paul and Juma, 1980) and are summarized in Table 1.

The concept of an active and a passive organic N phase (Jansson, 1958) was incorporated into the model. The active organic N phase was divided into three components: biomass, active-(C+N) and metabolite-(C+N). This was possible as the biomass and active N can be

Table 1. Initial values of N pools and techniques used for pool size determination.

| Pool | ^{14}N | ^{15}N | Technique |
|---------------------------------|---------------------------|-----------------|--|
| | (ug g ⁻¹ soil) | | |
| Biomass | 167 | 0.072 | Fumigation incubation |
| Active fraction and metabolites | 389 | 0.169 | Isotopic dilution and curve peeling |
| Old fraction | 2073 | 0 | Associated with old carbon (carbon dating) |
| Stabilized fraction | 1496 | 0 | By difference |

measured using the techniques summarized in Table 1. The passive organic N phase was divided into two components: stabilized N with a half life of ~35 years and old N with a half life of ~600 years. The dynamics of the mineralization processes are presented below:

a. *Microbial growth and N immobilization*

The dynamics of microbial growth and N immobilization were described by first-order kinetics. In most cases, the decomposition of C substrates is independent of microbial biomass size. In the model, the decomposition of C from any pool occurred if C was present in that pool and mineral N was not equal to zero. Thus,

$$dCX/dt = KX \cdot CX \cdot WCOF \cdot TCOF \quad (1)$$

where

dCX/dt = the rate of substrate decomposition (mg g⁻¹ soil d⁻¹)

KX = the decomposition rate constant for substrate CX (d⁻¹)

CX = the substrate C pool size (mg C g⁻¹ soil)

$WCOF$ and $TCOF$ = moisture and temperature reduction factors for sub-optimal conditions.

The values of these reduction factors ranged from 0 to 1 depending upon the physical conditions prevailing in soil (Van Veen, 1977).

The rate of C incorporation into the microbial biomass ($CXINC$; mg C g⁻¹ soil d⁻¹) was

$$CXINC = (YCOFX/100) \cdot dCX/dt \quad (2)$$

where $YCOFX$ is the efficiency of utilization of CX. The rate of CO₂-C evolution ($CXCO_2$, mg C g⁻¹ soil d⁻¹) during the decomposition of CX was:

$$CXCO_2 = (1 - (YCOFX/100)) \cdot dCX/dt \quad (3)$$

The rate of N incorporation into the microbial biomass ($CNINC$, mg N g⁻¹ soil d⁻¹) was

$$CNINC = (dCX/dt)/CNB \quad (4)$$

where CNB is the C:N ratio of the microbial biomass. Thus, the rate of N uptake during microbial growth was equal to the gross immobilization rate.

b. *Microbial biomass decay*

The decay of the biomass was described by first-order kinetics, i.e. the amount of biomass that decayed per unit time was proportional to its size. The decay C and N products were transferred into the active (C+N) and metabolite pools. The metabolites were divided into two classes: those containing C and N (proteins, amino sugars) and those containing C only (polysaccharides).

In the model, 50% of C and N of the microbial decay products were channelled into the active (C+N) pool. The remainder of the N was channelled into the metabolite (C+N) pool. The amounts of the C entering the metabolite (C+N) and metabolite C pools were calculated as follows:

Assuming a C:N ratio of three, the amount of C transferred to metabolite (C+N) pool was three times the amount of N transferred into it. The remainder of the C [50% of the total microbial C decay products minus C transferred into metabolite (C+N)] was channelled into metabolite-C pool.

c. *Mineralization of N*

Mineralization of N occurred during the decomposition of substrates containing C and N. The rate of N mineralization was calculated by multiplying the rate of C decomposition by the C:N ratio of the pool. Alternately, the rate of N mineralization was calculated by multiplying the decay rate constant (d^{-1}) for the (C+N) pool by the N content of the pool ($mg\ N\ g^{-1}\ soil$). The gross N mineralization rate was equal to the sum of the rates of N mineralization from various (C+N) pools.

d. *Stabilization of C and N*

Part of the active-(C+N) is transformed physically or chemically into the stabilized-(C+N). Further transformation of the stabilized-(C+N) results in the formation of recalcitrant organic matter. The amounts transferred from one pool to another were described by first-order kinetics.

8. INPUT PARAMETERS

The initial values, decay and transfer rate constants, efficiency of utilization of C substrates and C:N ratios of various (C+N) integrals needed in the model are summarized in Table 2.

Temperature and moisture content were time-dependent variables.

9. OUTPUT VARIABLES

The basic version of TRAMIN evaluates the following variables that are experimentally verifiable: biomass C and N, mineral N and CO_2 -C released during incubation. When tracer ^{15}N is used, the following additional variables are verifiable: biomass ^{15}N , mineral ^{15}N ,

Table 2. Initial values, decay and transfer rate constants, efficiency of utilization of C substrates and C:N ratios of various (C+N) integrals used in the model.

| Pool | C | ¹⁴ N | ¹⁵ N | Decay rate constant | Efficiency of utilization of C | C:N ratio |
|-------------------------------|------------------------------|----------------------------|----------------------------|---------------------|--------------------------------|-----------|
| | ($\mu\text{g g}^{-1}$ soil) | (ng g^{-1} soil) | (ng g^{-1} soil) | (d^{-1}) | (%) | |
| Biomass | 1,000 | 167 | 72 | 0.0143 | - | 6 |
| Active fraction ¹ | 2,302 | 384 | 167 | 0.0037 | 40 | 6 |
| Metabolite-C | 15 | - | - | 1.0 | 60 | - |
| Metabolite-(C+N) ² | 15 | 5 | 2 | 1.0 | 60 | 3 |
| Stabilized-(C+N) ² | 17,542 | 1,496 | 0 | $6.0 \cdot 10^{-4}$ | 40 | 11 |
| Old-(C+N) | 24,455 | 2,073 | 0 | $3.0 \cdot 10^{-6}$ | - | 11.8 |
| Decomposable C | 2 | - | - | 1.0 | 60 | - |
| Slowly decomposable C | 50 | - | - | 0.1 | 60 | - |
| Mineral N | - | 21 | 9 | - | - | - |

1. Transfer rate constant of active material to stabilized fraction = $0.5 \cdot 10^{-3}$ (d^{-1}).

2. Transfer rate constant of stabilized fraction to old fraction = $3.0 \cdot 10^{-6}$ (d^{-1}).

active ¹⁵N and active ¹⁴N. The model calculates the size of N integrals and their atom % abundances.

When ¹⁴C is used, the following variables are verifiable: biomass ¹⁴C, ¹⁴CO₂-C, active ¹⁴C (sum of active (C+N), metabolite C and metabolite (C+N)). The model calculates the size of C integrals and their specific activities.

When ¹⁴C and ¹⁵N are used all the above mentioned variables are verifiable.

9. OBSERVATIONS

Enriched ¹⁵N soil samples were obtained from a field experiment conducted on Weirdale loam, a Gray-Black Chernozemic soil. These surface samples were incubated in the laboratory for 12 weeks at field capacity moisture content and 28±1 °C. During the incubation, total N, biomass N, active N and mineral N and their atom % ¹⁵N abundances were determined. Also, biomass-C and CO₂-C evolved were measured. Detailed experimental procedures are summarized elsewhere (Paul and Juma, 1980).

The TRAMIN programme was used to simulate the ¹⁴N, ¹⁵N and C transformations.

11. COMPARISON OF RESULTS

The predicted values of total organic ¹⁵N remaining in soil and cumulative CO₂-C evolved were slightly higher than the experimental values (Table 3); the mineral ¹⁵N level predicted was lower than the actual value. The other outputs were similar indicating that the gross, transfer rate constants and pool sizes used (Table 2) were of proper magnitude.

The dynamics of various ¹⁴N and ¹⁵N pools as predicted by the model are shown in Table 4. During the 12 weeks incubation, when no exogenous C was supplied, the biomass with an initial size of 72 (ng g^{-1}) immobilized 56 $\text{ng } ^{15}\text{N g}^{-1}$ soil and decayed by 76 ng^{-1} resulting in a final size of 53 ng g^{-1} (Fig. 2). The metabolite N mineralized rapidly and constituted

Table 3. Comparison of experimental data with simulation model outputs at the end of 12 weeks incubation.

| Pool | Experimental | Simulated |
|---|---------------------------|-----------|
| | (µg g ⁻¹ soil) | |
| Total organic ¹⁵ N remaining in soil | 0.203 | 0.211 |
| Biomass- ¹⁵ N | 0.052 | 0.053 |
| Active fraction + metabolite- ¹⁵ N | 0.152 | 0.150 |
| Stabilized- ¹⁵ N fraction | - | 0.007 |
| Mineral- ¹⁵ N | 0.045 | 0.039 |
| Mineral- ¹⁴ N | 115 | 121 |
| Cumulative CO ₂ -C evolved | 938 | 1153 |

Table 4. Dynamics of various N fractions in soil during 12 weeks incubation.

| Fraction | Simulated values | | | | Experimental value at end of 12 weeks |
|--|------------------|--------|---------|-------------------------|---------------------------------------|
| | Initial size | Inputs | Outputs | Size at end of 12 weeks | |
| <i>¹⁵N (ng g⁻¹ soil)</i> | | | | | |
| Biomass-N | 72 | 56 | 76 | 53 | 52 |
| Metabolite-N | 2 | 38 | 39 | 1 | 152 |
| Active-N | 166 | 38 | 54 | 150 | - |
| Stabilized-N | 0 | 7 | 0.1 | 7 | - |
| Old-N | 0 | 0 | 0 | 0 | - |
| Mineral-N | 9 | 86 | 56 | 39 | 45 |
| <i>¹⁴N (µg g⁻¹ soil)</i> | | | | | |
| Biomass-N | 167 | 163 | 187 | 143 | 152 |
| Metabolite-N | 5 | 94 | 98 | 1 | - |
| Active-N | 384 | 94 | 125 | 352 | - |
| Stabilized-N | 1496 | 15 | 55 | 1456 | - |
| Old-N | 2073 | 0.1 | 0.4 | 2073 | - |
| Mineral-N | 21 | 263 | 163 | 121 | 115 |

only a small pool at any time. The active N with an initial size of 166 ng ¹⁵N g⁻¹ increased by 38 ng g⁻¹ and released 54 ng resulting in a final size of 150 ng g⁻¹. A small portion (7 ng g⁻¹) of the ¹⁵N released from active-N was transferred to the stabilized fraction while the rest was mineralized. At the end of the 12 weeks incubation, the ¹⁵N content of the stabilized-N showed a net increase while the ¹⁵N content of biomass and active N showed a net decrease. During the 12 weeks incubation, the model predicted a net mineralization of 30 ng g⁻¹ ¹⁵N. The relative net changes of the ¹⁵N pools expressed as a fraction of ¹⁵N mineralized were biomass (-63%), active-N (-53%), metabolite-N (-3%), stabilized-N (+23%), old-N (0%). At present, it is not possible to measure the ¹⁵N in stabilized-N and old-N.

During the 12 weeks incubation, the biomass with an initial size of 167 µg ¹⁴N g⁻¹ soil immobilized 163 µg g⁻¹ and decayed by 187 µg g⁻¹ resulting in a net decrease of 24 µg g⁻¹ (Fig. 3). The metabolite-N and active-N with equal inputs also declined. The stabilized-N showed a net mineralization of 40 µg g⁻¹ while the old N contributed 0.3 µg g⁻¹. Gross

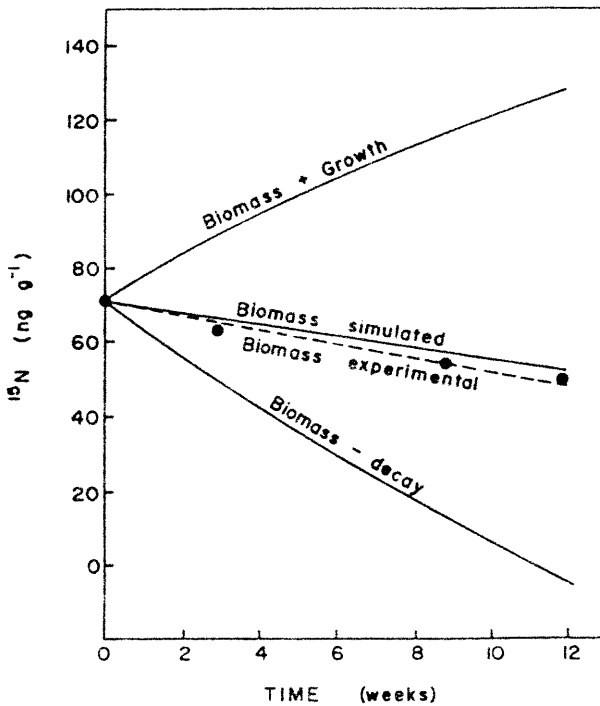


Fig. 2. Biomass-¹⁵N: cumulative growth and decay rates.

mineralization during this period amounted to $263 \mu\text{g g}^{-1}$, while the net mineralization was $100 \mu\text{g N g}^{-1}$. The relative net contribution of N fractions to mineral ¹⁴N were biomass (74%), metabolites (4%), active-N (32%), stabilized-N (40%), old-N (0.3%). At present, it is not possible to determine the sizes of metabolite-N, active-N, stabilized-N and old-N.

The turnover time of any pool can be calculated from the tracer data provided the system is in steady state conditions. The soil system generally is not in a steady state condition on a short term basis. Therefore, it is necessary to use simulation techniques to get a good estimate of the fluxes in and out of various pools. In the present version of the TRAMIN model, the C, ¹⁴N and ¹⁵N transformations were simulated. The major advantage of simulating the N isotopes separately is that it helps to quantify these fluxes accurately. For any pool, the ¹⁴N is a good indication of changes in size while the ¹⁵N is an excellent index of the activity within the pool. Sensitivity analysis showed that the decay or transfer rate constants used in the model had a very narrow range of values.

Jansson (1958) showed that mineralization and immobilization of N were continuous processes and demonstrated the internal cycling of N. The model predicted that ¹⁵N in soil undergoes rapid turnover. During the 12 weeks incubation, 36% (86 ng g^{-1}) of the ¹⁵N initially present in soil was mineralized while 23% (56 ng g^{-1}) was reimmobilized. The net change of 30 ng g^{-1} was ~13% of the total ¹⁵N at the beginning of the experiment. Thus,

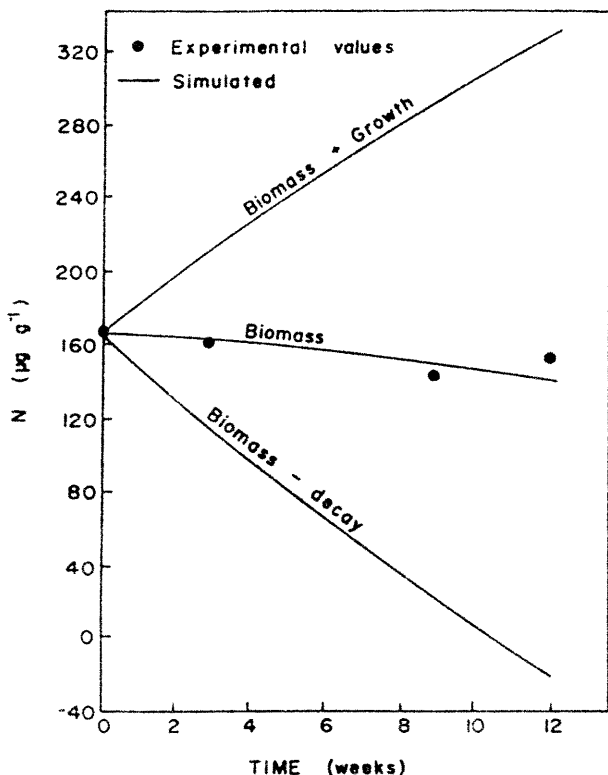


Fig. 3. Total microbial biomass-N: cumulative growth and decay rates.

net mineralization was 1/3 of the gross mineralization, i.e. of every 3 units of N that were mineralized, 2 were reimmobilized, resulting in a net mineralization of 1 unit. Therefore, the half life of ¹⁵N in soil organic matter is very long due to the internal cycling of N.

Jansson (1958) suggested that the active organic N phase was the major source of mineralized N. Stanford and Smith (1972) used long term incubations and mathematical techniques to estimate the size and decay rate constant of the mineralizable N pool for 39 soils which were in net mineralization conditions. Thus, the decay rate constant they obtained was related to net mineralization only. The TRAMIN model after considering gross mineralization and immobilization rates predicted that all the N fractions in soil contributed to the mineral N. In this study the active organic N phase comprised of the biomass, active-N and metabolic-N and accounted for almost all of the ¹⁵N mineralized, but contributed ~60% of the ¹⁴N mineralized. The stabilized-N and old-N accounted for the rest of ¹⁴N mineralized. The model predicted that various organic N pools are mineralizable and have different mineralization rate constants. Therefore, it is difficult to describe the soil processes involved in N mineralization by a single first-order decay equation, es-

pecially when a net mineralization rate constant is used in the equation.

12. LIMITS AND LIMITATIONS

The model assumes aerobic conditions at all the times and does not consider nitrification or N losses and gains. The N in soil was divided into biomass, active-N, metabolite-N, stabilized-N and old-N, although it is well known that the N compounds in soil form a continuum from simple soluble to recalcitrant compounds. These compounds were grouped into the above mentioned pools based on the decay rates of the pools as determined by tracer and isotopic dilution techniques. At present, it is not possible to measure the size and decay rate constants of the stabilized-N and old-N, however, a wealth of information exists in the literature which supports the concept that large pools or fractions with long half lives exist (Oades and Ladd, 1977).

The mineralization process described in the model assumes that when a substrate containing C and N is decomposed, the N is first converted to NH_4^+ . Thus, compounds like simple amino acids and amino sugars are deaminated before being taken up by the biomass. The amount of N immobilized is governed by the amount of C incorporated into the biomass and by the C:N ratio of the biomass.

13. COMPUTER

IBM 360/370.

14. PROGRAMME LANGUAGE

CSMP III.

15. RUNNING TIME/COST

1.5 CPU minutes, cost \$16.00 for simulating 84-day laboratory experiment.

USERS

The programme has been recently developed, however, the concepts developed could be applied to other soil systems where the role of micro-organisms is explicitly defined in the mineralization and immobilization of N.

17. DEVELOPERS AND PRINCIPAL CONTACT

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